



Hydrologidag, 2014

# Application of satellite data to estimate evapotranspiration over large areas.

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# Introduction

## 1. What is remote sensing?

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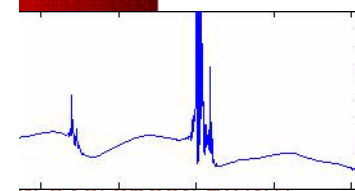
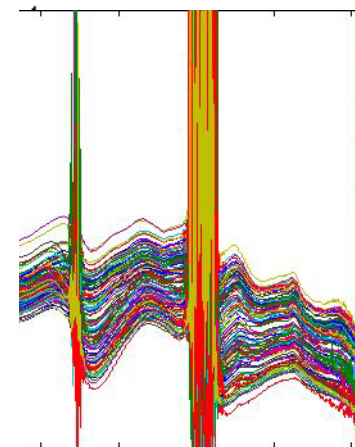
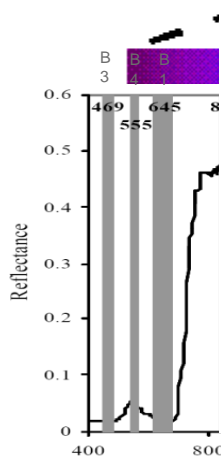
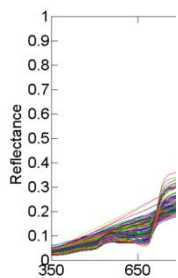
Table 1: List of indices calculated in this study using MODIS and the field spectra measurements.

Index	Formula	Reference
Normalized difference vegetation index	$NDVI = \frac{B_2 - B_1}{B_2 + B_1}$	(Tucker, 1979)
Enhanced Vegetation Index	$EVI = 2.5 \cdot \left( \frac{B_2 - B_1}{B_2 + 6 \cdot B_1 - 7.5 \cdot B_5} \right)$	(Huete et al., 2002)
Normalized Difference Water Index	$NDWI = \frac{B_2 - B_5}{B_2 + B_5}$	(Gao, 1996)
Short Infrared Water Stress Index	$SIWSI = \frac{B_6 - B_2}{B_6 + B_2}$	(Fensholt & Sandholt, 2003)
Normalized Difference Water Index	$NDWI = \frac{B_3 - B_7}{B_3 + B_7}$	(Rubio et al., 2006)
Normalized Difference Infrared Index	$NDII = \frac{B_7 - B_6}{B_7 + B_6}$	(Hardisky et al., 1983)
Simple Ratio Water Index	$SRWI = \frac{B_2}{B_5}$	(Zarco-Tejada et al., 2003)
Inverted Simple Ratio Water Index	$ISRWI = \frac{B_5}{B_2}$	(Trombetti et al., 2008)
Moisture Stress Index	$MSI = \frac{B_6}{B_2}$	(Jr. & Rock, 1989)
Moisture Stress Index 7	$MSI7 = \frac{B_7}{B_2}$	(Trombetti et al., 2008)
Inverted Moisture Stress Index	$IMSI = \frac{B_2}{B_6}$	This study
Inverted Moisture Stress Index 7	$IMSI7 = \frac{B_2}{B_7}$	This study
Shortwave Infrared Ratio	$SRI = \frac{B_6}{B_7}$	(Trombetti et al., 2008)
Soil Adjusted Vegetation Index	$SAVI = \left( \frac{B_2 - B_1}{B_2 + B_1 + L} \right) \cdot (1 + L)$	(Huete, 1988)
Global Environment Monitoring Index	$GEMI = \eta \cdot (1 - 0.25\eta) - \frac{\rho_1 - 0.125}{1 - \rho_1}$ where $\eta = \frac{2 \cdot (\rho_2^2 - \rho_1^2) + 1.5 \cdot \rho_2 + 0.5 \cdot \rho_1}{\rho_2 + \rho_1 + 0.5}$	(Pinty & Verstraete, 1992)
Canopy Water Content	See Trombetti et al. 2008	(Trombetti et al., 2008)
Visible Atmospherically Resistant Index	$VARI = \frac{B_4 - B_1}{B_4 + B_1 - B_3}$	(Gitelson et al., 2002a)
Global Vegetation Monitorin Index	$GVMI = \frac{(NIR_{REC} + 0.1) - (SWIR_{REC} - 0.02)}{(NIR_{REC} + 0.1) + (SWIR_{REC} - 0.02)}$	(Ceccato et al., 2002)
Normalized indices calculated in the study	$NormalizedIndex = \frac{B_x - B_y}{B_x + B_y}$ where $B_x$ and $B_y$ indicate one of the seven reflectance bands available in the MOD09GA product	

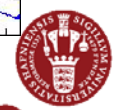
out

## 2. Typ

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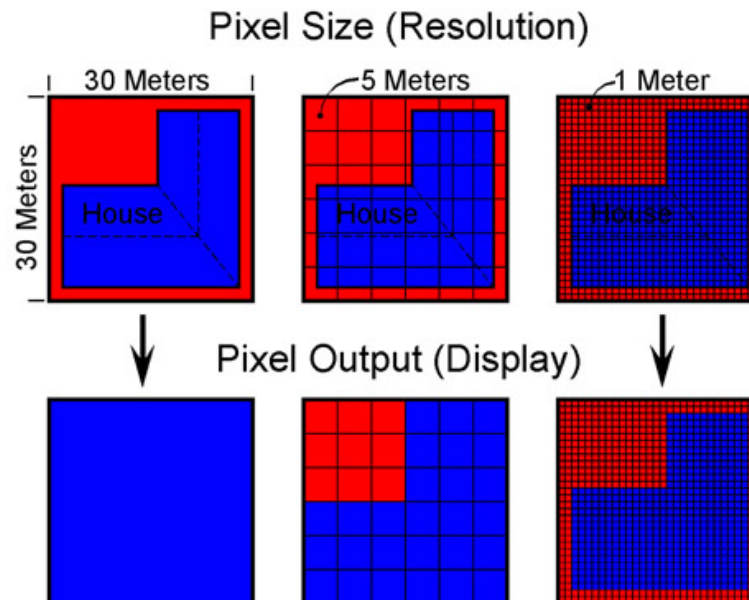


wavelength



## Introduction

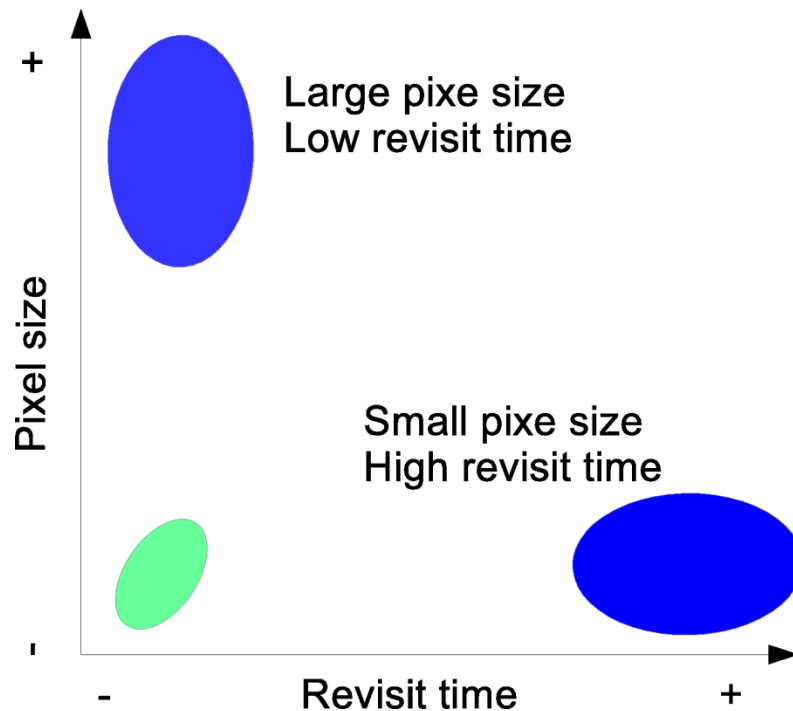
1. What is remote sensing?
  - Observation and measurement of an object without touching it.
2. Types of resolution in remote sensing
  - Spatial Resolution



Source: <http://www.satimagingcorp.com/services/resources/characterization-of-satellite-remote-sensing-systems/>

## Introduction

1. What is remote sensing?
  - Observation and measurement of an object without touching it.
2. Types of resolution in remote sensing
  - Temporal Resolution  
How often the data is collected



## Introduction

- Sources of remote sensing data

### UAV

- Small pixel size (cm)

- Difficulties to cover a large area
- Limited sensors can be mounted

### Airborne

- Small pixel size (m)
- Hyperspectral sensors can be mounted

- Difficulties to cover a large area
- Expensive
- Require a lot of planification

### Polar orbiting

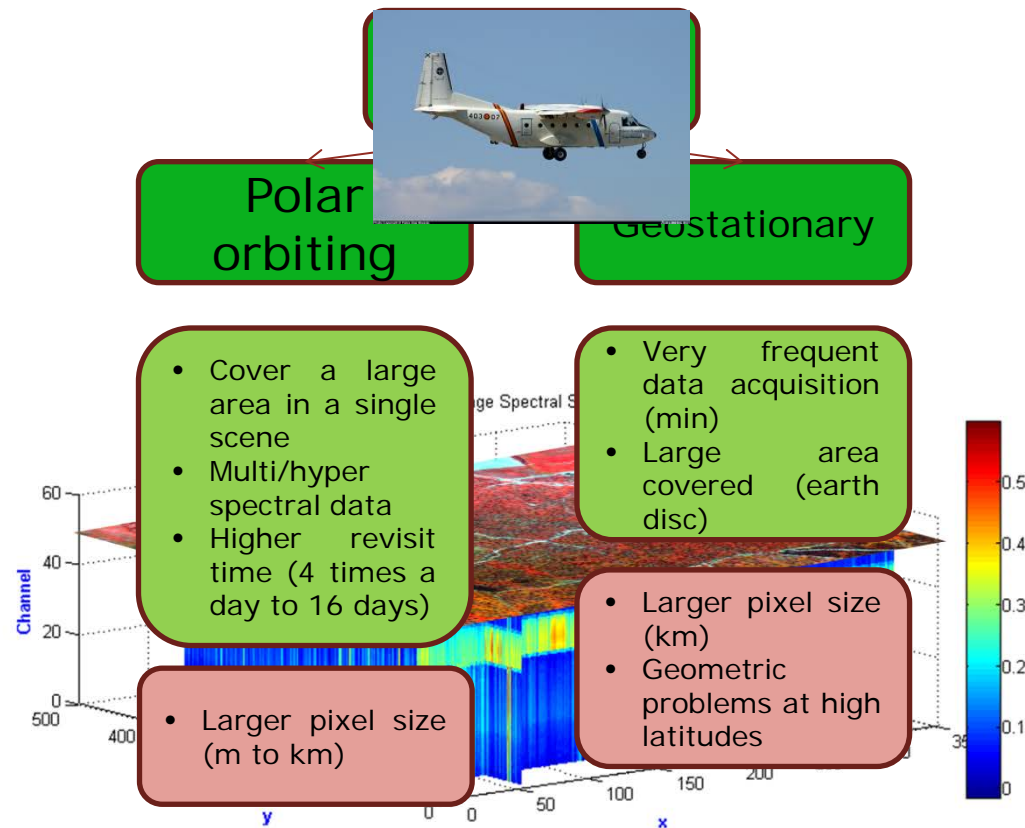
- Cover a large area in a single scene
- Multi/hyper spectral data
- Higher revisit time (4 times a day to 16 days)

- Larger pixel size (m to km)

### Geostationary

- Very frequent data acquisition (min)
- Large area covered (earth disc)

- Larger pixel size (km)
- Geometric problems at high latitudes



# How do we calculate Evapotranspiration (ET) using satellites?

- Empirical approach
  - Vegetation index related with ET measurements
- Semi-empirical approach
  - Triangle approach relating temperature and vegetation indices
- Physical approaches (Based on Land Surface Temperature (LST), Leaf Area Index and Albedo)
  - One source energy balance model (OSEB)
  - Two source energy balance model (TSEB)
    - Dual-Temperature-Difference (DTD) adapted for polar orbiting satellites (Guzinski, 2013)**



## The DTD model

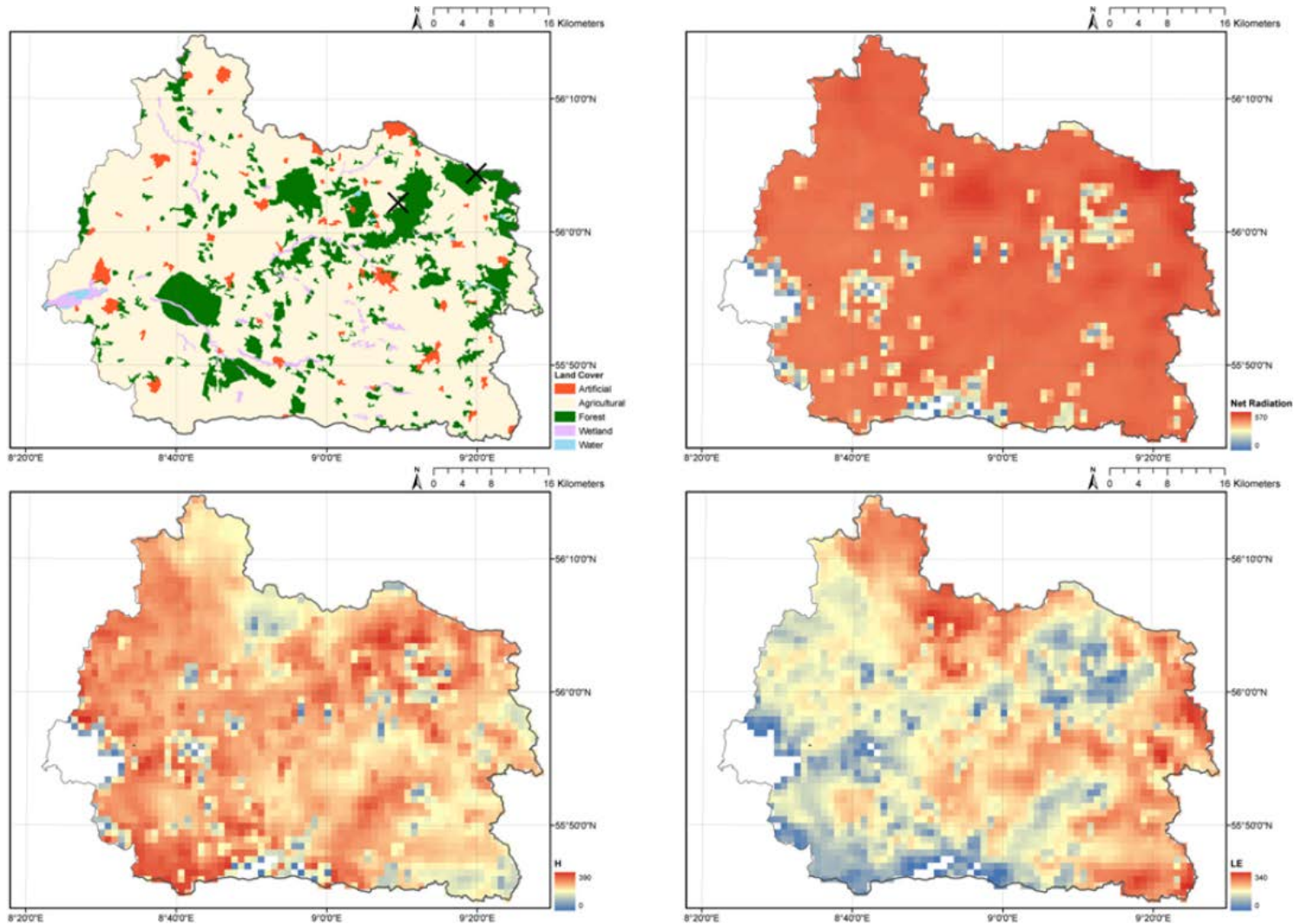
- Is based on the energy conservation law

$$R_n = LE + H + G$$

- Uses two observations in time to minimize systematic errors in temperature estimation
- Requires LAI as input to split LST between soil and canopy
- Net radiation and ET for soil and canopy as outputs.



## The DTD applied to Skjern River catchment



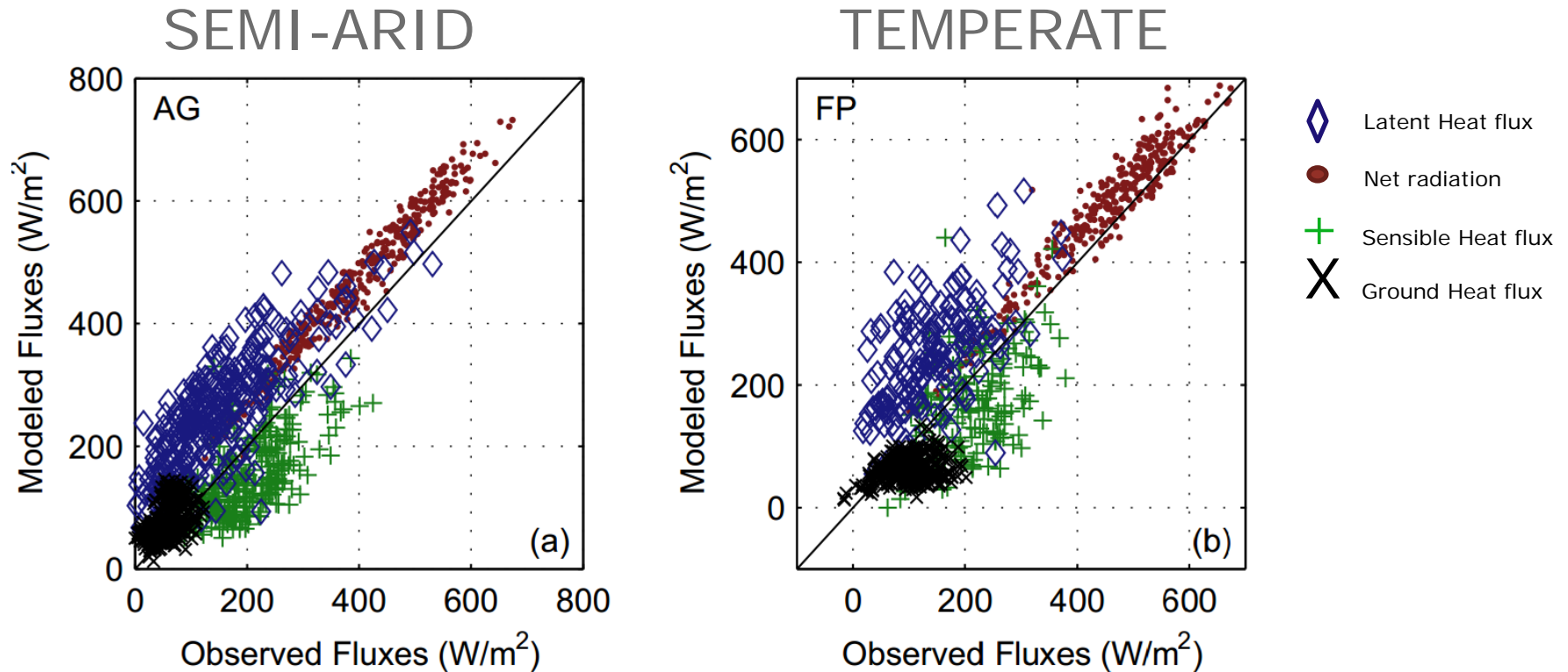
Guzinski et al. (2013)

Hydrologidag, Odense. October 2014





## DTD performance

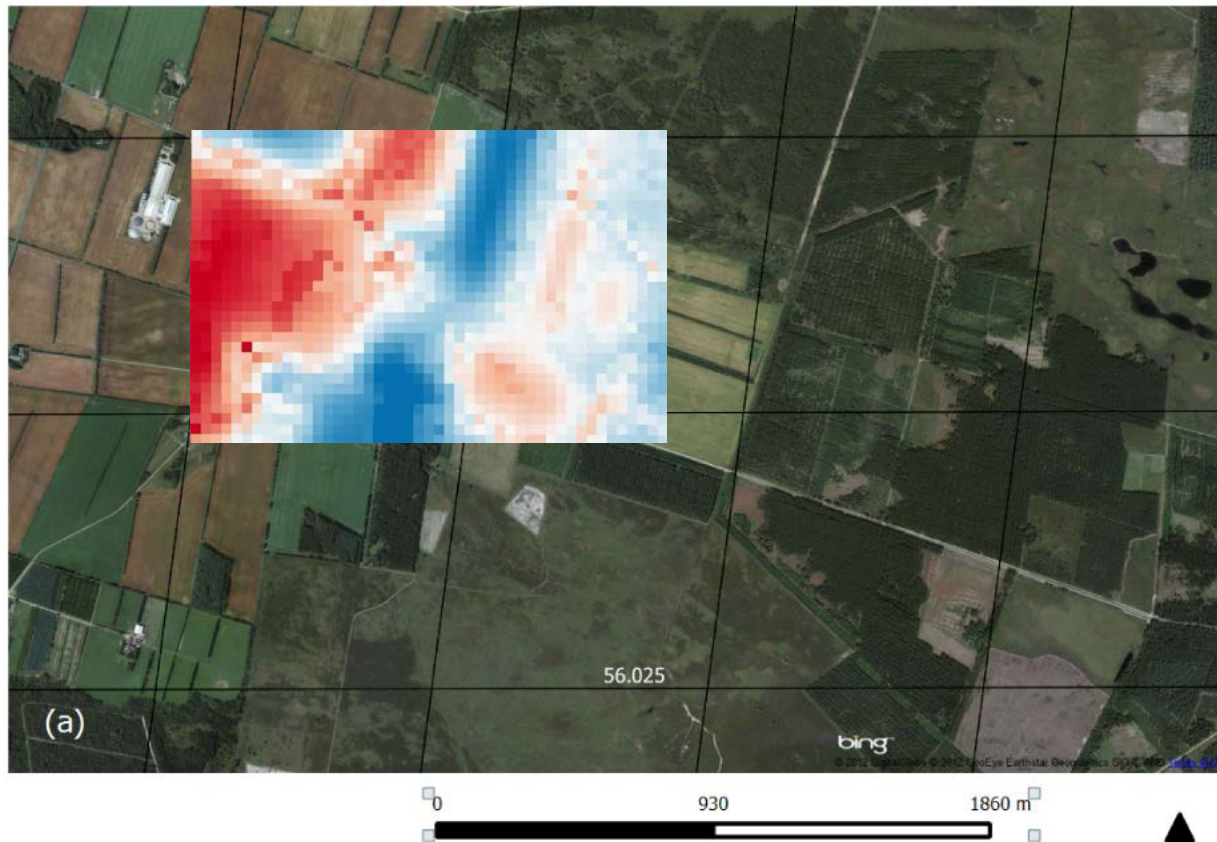


(Guzinski, 2013)



## Scale issues

- Downscaling fluxes
  - Low resolution fluxes can be disaggregated using higher spatial resolution images (Dis-ALEXI)



## Summary and conclusions

- Remote sensing is a powerful tool to monitor hydrological processes
- ET was presented as an example. Spectral and thermal information can also be used to retrieve other biophysical properties of vegetation (LAI, Albedo) as well as other soil properties
- Combination of different sensors offers a wide range of possibilities, as very dynamic processes can be monitored with higher spatial resolution. Specially important when studying heterogeneous areas.



Thanks for your attention!!!!

Any questions???

